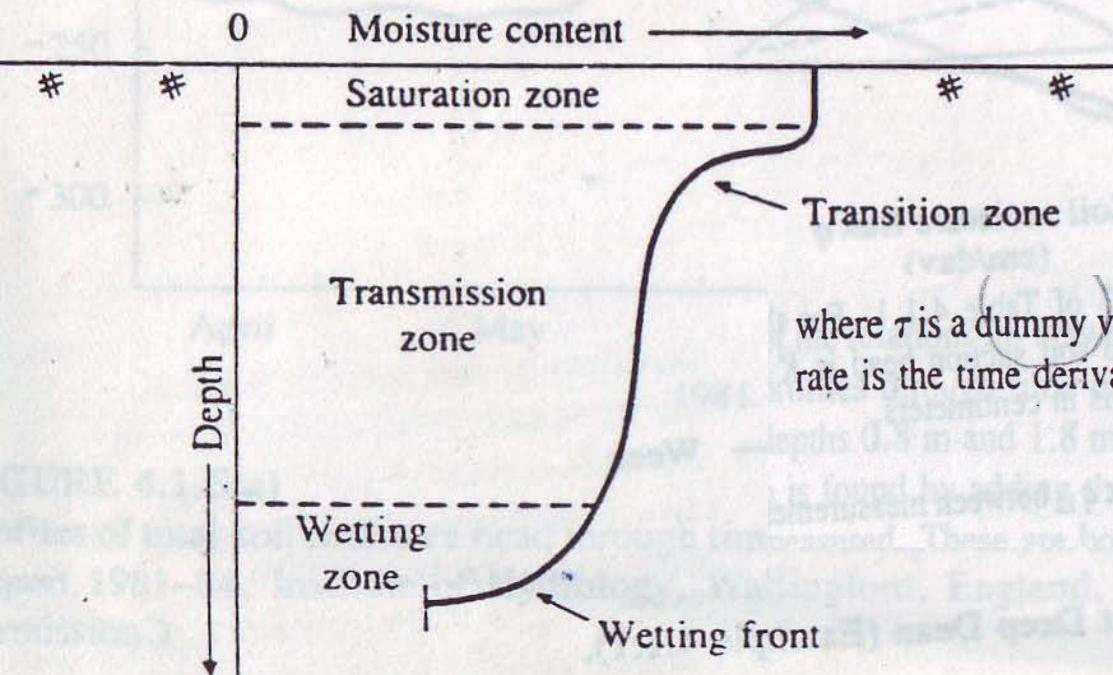


INFILTRASI

Konsep Infiltrasi

$$F(t) = \int_0^t f(\tau) d\tau \quad (4.2.1)$$



where τ is a dummy variable of time in the integration. Conversely, the infiltration rate is the time derivative of the cumulative infiltration:

$$f(t) = \frac{dF(t)}{dt}$$

FIGURE 4.2.1
Moisture zones during infiltration.

Horton's Equation

One of the earliest infiltration equations was developed by Horton (1933) who observed that infiltration begins at some rate f_0 and exponentially decreases until it reaches a constant rate f_c (Fig. 4.2.2):

$$f(t) = f_c + (f_0 - f_c)e^{-kt}$$

where k is a decay constant having dimensions $[T^{-1}]$. Eagleson (1962) and Raudkivi (1979) have shown that Horton's equation can be derived from the equation (4.1.12) by assuming that K and D are constants independent of moisture content of the soil. Under these conditions (4.1.12) reduces to

$$\frac{\partial \theta}{\partial t} = D \frac{\partial^2 \theta}{\partial z^2}$$

which is the standard form of a diffusion equation and may be solved for the moisture content θ as a function of time and depth. Horton's equation can be obtained from solving for the rate of moisture diffusion $D(\partial \theta / \partial z)$ at the soil surface.

Penyelesaian Metoda Horton

- Data di ambil dari hasil pengamatan lapangan
- Alat yang digunakan Double Infiltrometer

atau

$$f(t) = f_c + (f_o - f_c)e^{-kt}$$

$$f(t) - f_c = (f_o - f_c)e^{-kt}$$

$$\frac{f(t) - f_c}{f_o - f_c} = e^{-kt}$$

$$\ln(e^{-kt}) = \ln\left(\frac{f(t) - f_c}{f_o - f_c}\right)$$

$$-kt = \ln\left(\frac{f(t) - f_c}{f_o - f_c}\right)$$

$$k = -\frac{\ln\left(\frac{f(t) - f_c}{f_o - f_c}\right)}{t}$$

$$f(t) = f_c + (f_o - f_c)e^{-kt}$$

$$f(t) - f_c = (f_o - f_c)e^{-kt}$$

$$\ln(f(t) - f_c) = \ln(f_o - f_c) + \ln(e^{-kt})$$

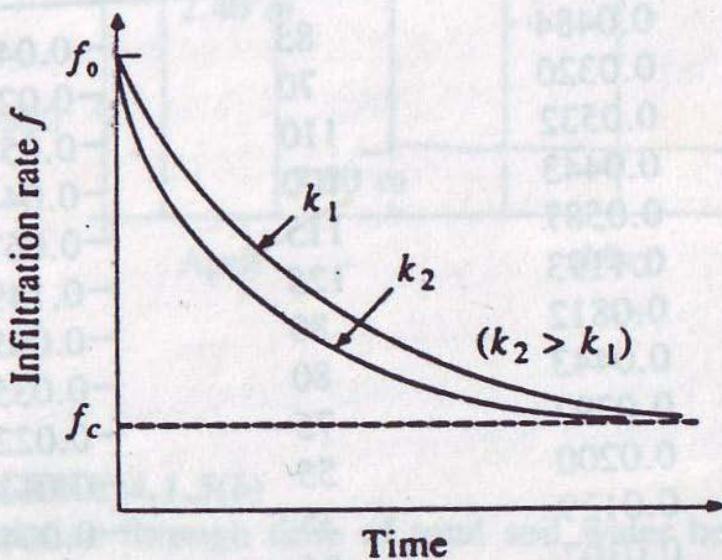
$$\ln(f(t) - f_c) - \ln(f_o - f_c) = -kt \cdot \ln e$$

$$-kt = \ln(f(t) - f_c) - \ln(f_o - f_c)$$

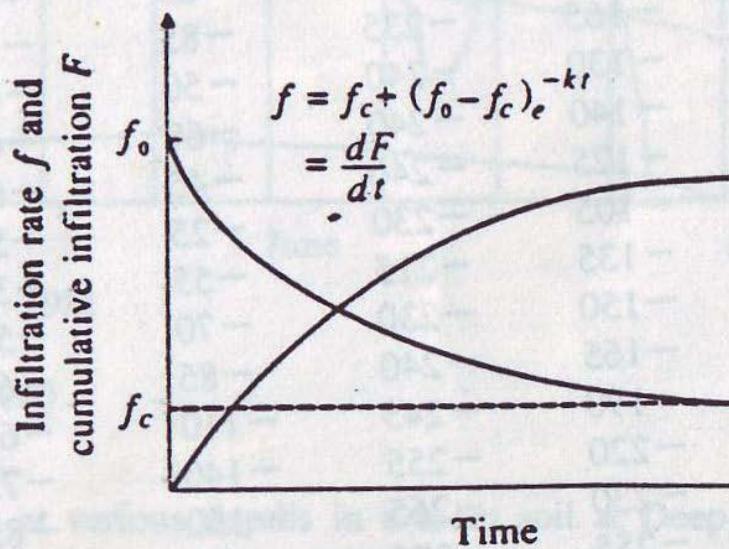
$$k = \frac{\ln(f_o - f_c) - \ln(f(t) - f_c)}{t}$$

Phillip's Equation

Philip (1957, 1969) solved Richard's equation under less restrictive conditions by assuming that K and D can vary with the moisture content θ . Philip employed the Boltzmann transformation $B(\theta) = zt^{-1/2}$ to convert (4.1.12) into an ordinary differential equation in B , and solved this equation to yield an infinite series solution.



(a) Variation of the parameter k .



(b) Infiltration rate and cumulative infiltration.

FIGURE 4.2.2

Infiltration by Horton's equation.

lative infiltration $F(t)$, which is approximated by

$$F(t) = St^{1/2} + Kt \quad (4.2.5)$$

• S is a parameter called *sorptivity*, which is a function of the soil suction
tial, and K is the hydraulic conductivity.

By differentiation

$$f(t) = \frac{1}{2}St^{-1/2} + K \quad (4.2.6)$$

→ ∞, $f(t)$ tends to K . The two terms in Philip's equation represent the effects
il suction head and gravity head, respectively. For a horizontal column of
soil suction is the only force drawing water into the column, and Philip's
ion reduces to $F(t) = St^{1/2}$.

Example 4.2.1. A small tube with a cross-sectional area of 40 cm^2 is filled with soil and laid horizontally. The open end of the tube is saturated, and after 15 minutes, 100 cm^3 of water have infiltrated into the tube. If the saturated hydraulic conductivity of the soil is 0.4 cm/h , determine how much infiltration would have taken place in 30 minutes if the soil column had initially been placed upright with its upper surface saturated.

Solution. The cumulative infiltration depth in the horizontal column is $F = 100 \text{ cm}^3 / 40 \text{ cm}^2 = 2.5 \text{ cm}$. For horizontal infiltration, cumulative infiltration is a function of soil suction alone so that after $t = 15 \text{ min} = 0.25 \text{ h}$,

$$F(t) = St^{1/2}$$

and

$$2.5 = S(0.25)^{1/2}$$

$$S = 5 \text{ cm} \cdot \text{h}^{-1/2}$$

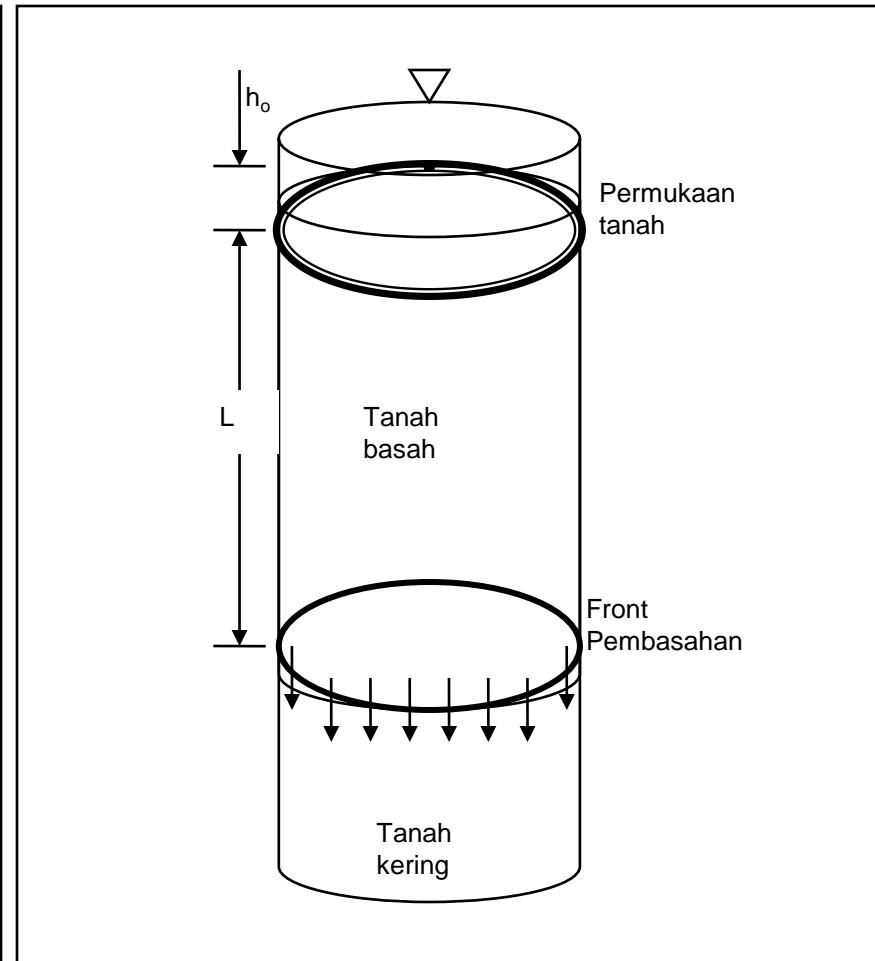
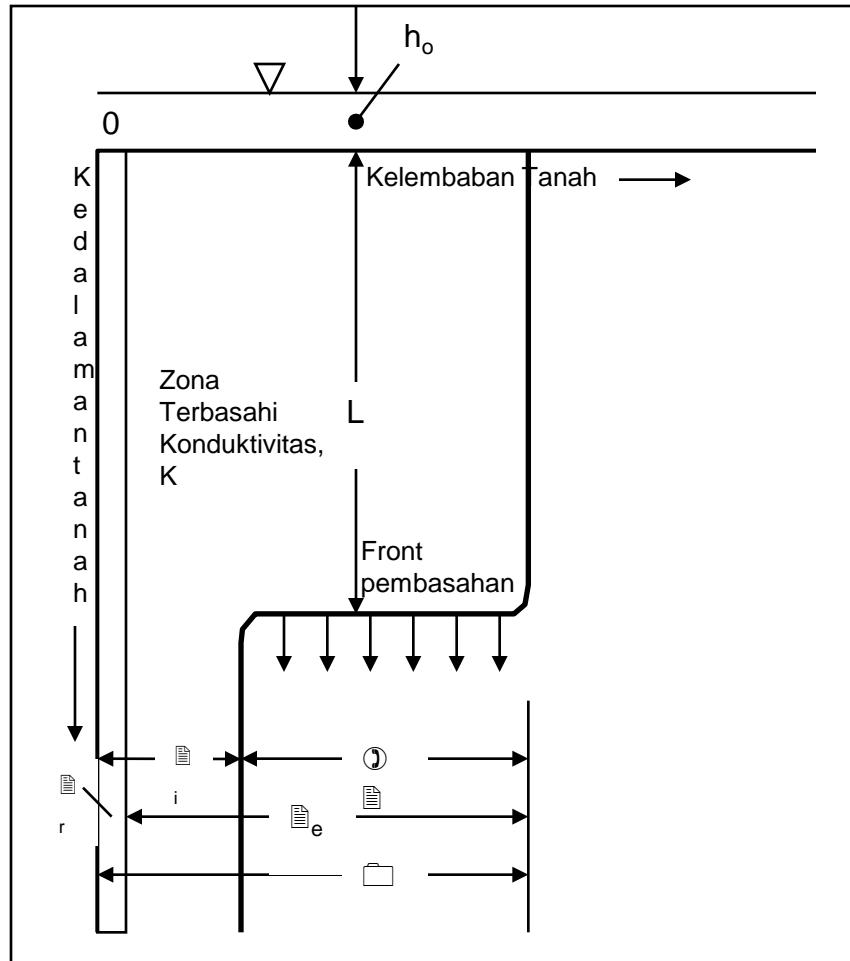
For infiltration down a vertical column, (4.2.5) applies with $K = 0.4 \text{ cm/h}$. Hence, with $t = 30 \text{ min} = 0.5 \text{ h}$

$$\begin{aligned} F(t) &= St^{1/2} + Kt \\ &= 5(0.5)^{1/2} + 0.4(0.5) \\ &= 3.74 \text{ cm} \end{aligned}$$

Double Infiltrometer untuk Pengukuran Infiltrasi



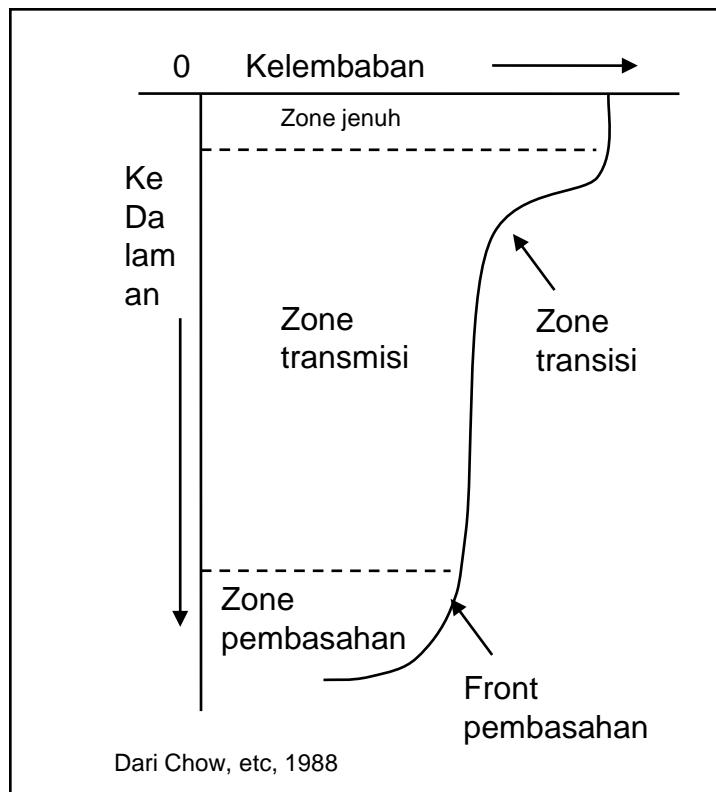
Teori Fisik : Metode Infiltrasi Green-Ampt



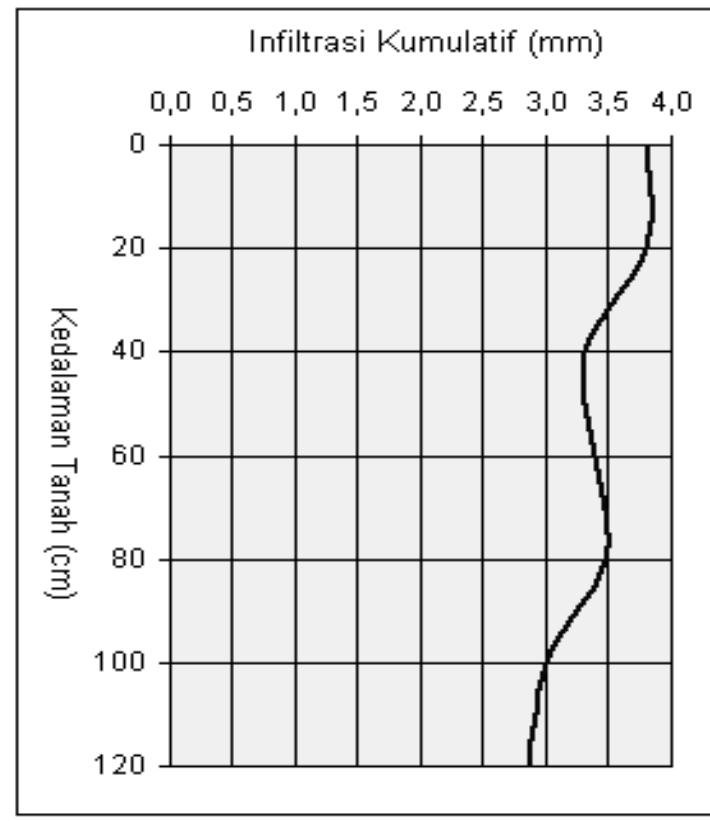
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Zonasi Kelembaban Tanah Menurut Kedalaman Tanah



Teoritik

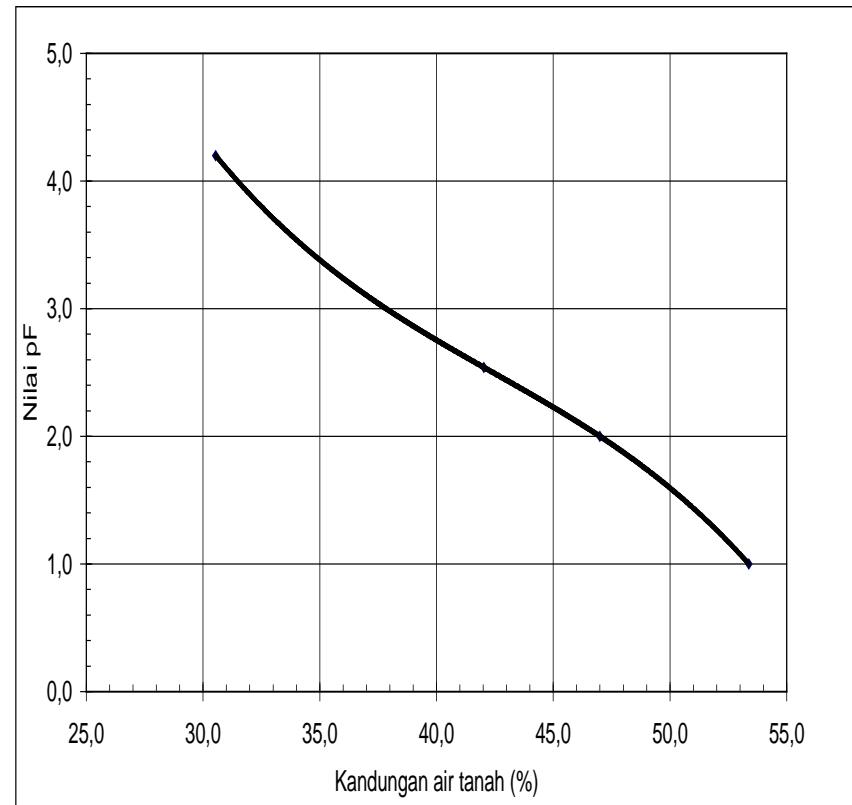


Empirik

Pendugaan Suction Head (✿) Berdasarkan Kelembaban Tanah

- $pF = \log (-\Psi)$, atau
- $pF = 29,30 - 1,684 \theta + 0,0371 \theta^2 - 0,00029 \theta^3$
- $\Psi = -(10^{pF})$
- $\theta = \text{kelembaban tanah (\%)}$

$$\psi = -(10)^{29,30 - 1,684\theta + 0,0371\theta^2 - 0,00029\theta^3}$$

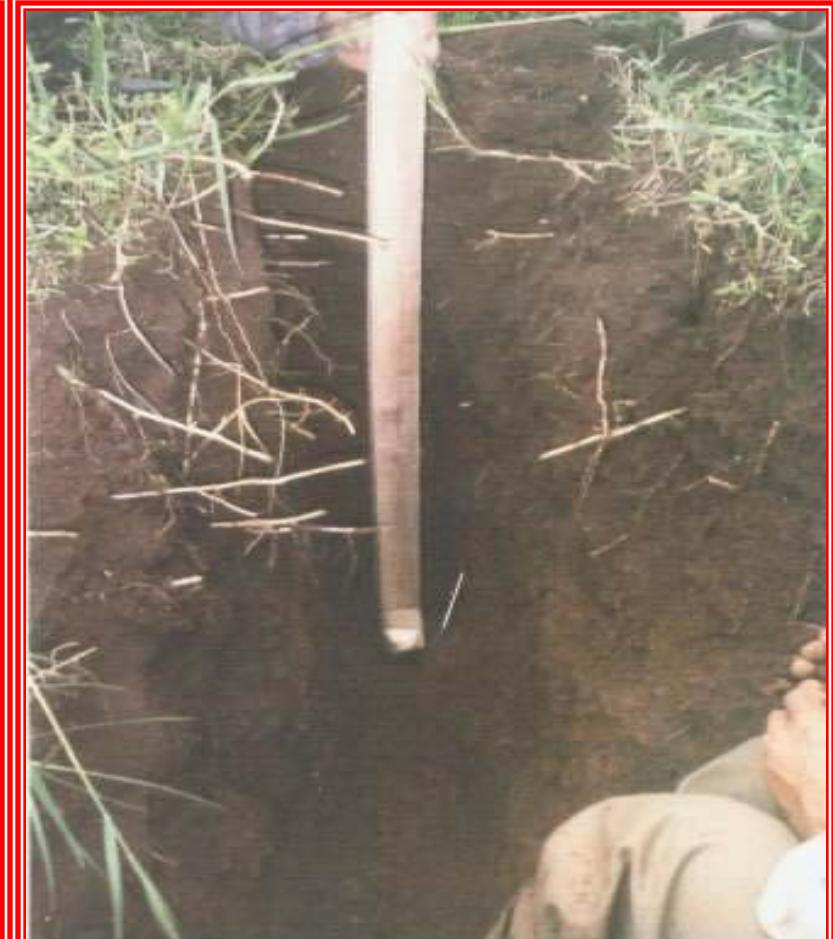


Pendugaan Nilai Permeabilitas (K) Berdasarkan Sifat Fisik Tanah

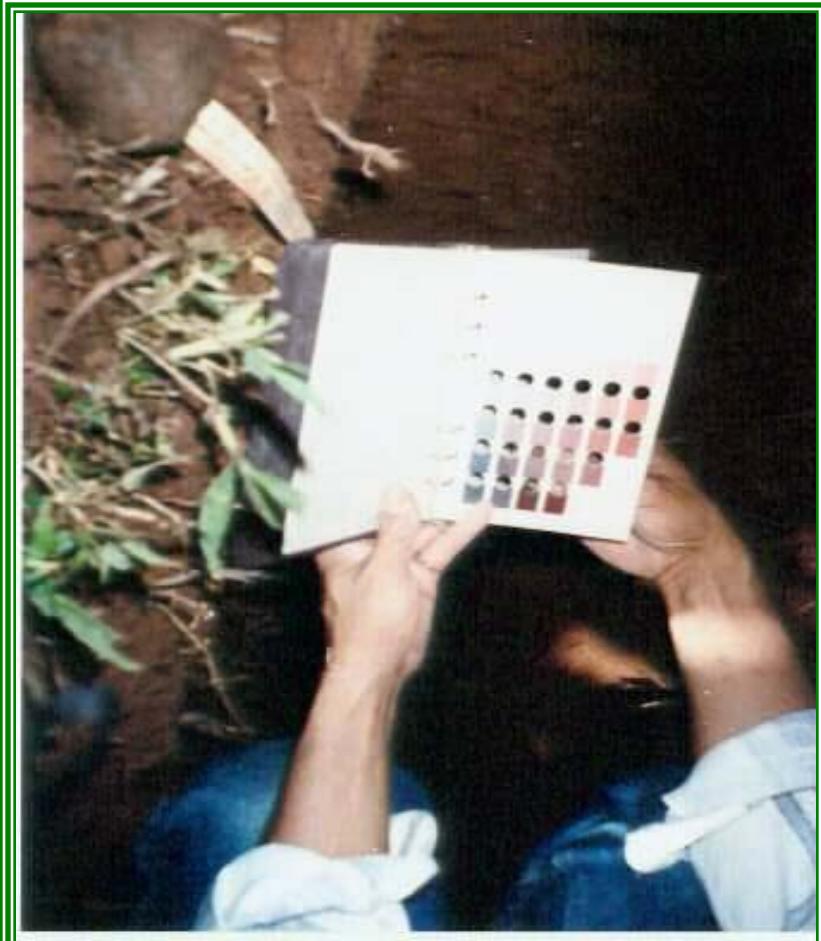
$$\ln K = -2,391 - 0,090.\theta + 0,161.\eta_c + 0,845.\eta_l$$

- $\theta : K = e^{-2,391-0,090.\theta+0,161.\eta_c+0,845.\eta_l}$
- η_c = kandungan pori drainase cepat (%)
- η_l = kandungan pori drainase lambat (%)

Pengambilan Sampel Tidak Terganggu untuk Uji Sifat Fisik Tanah



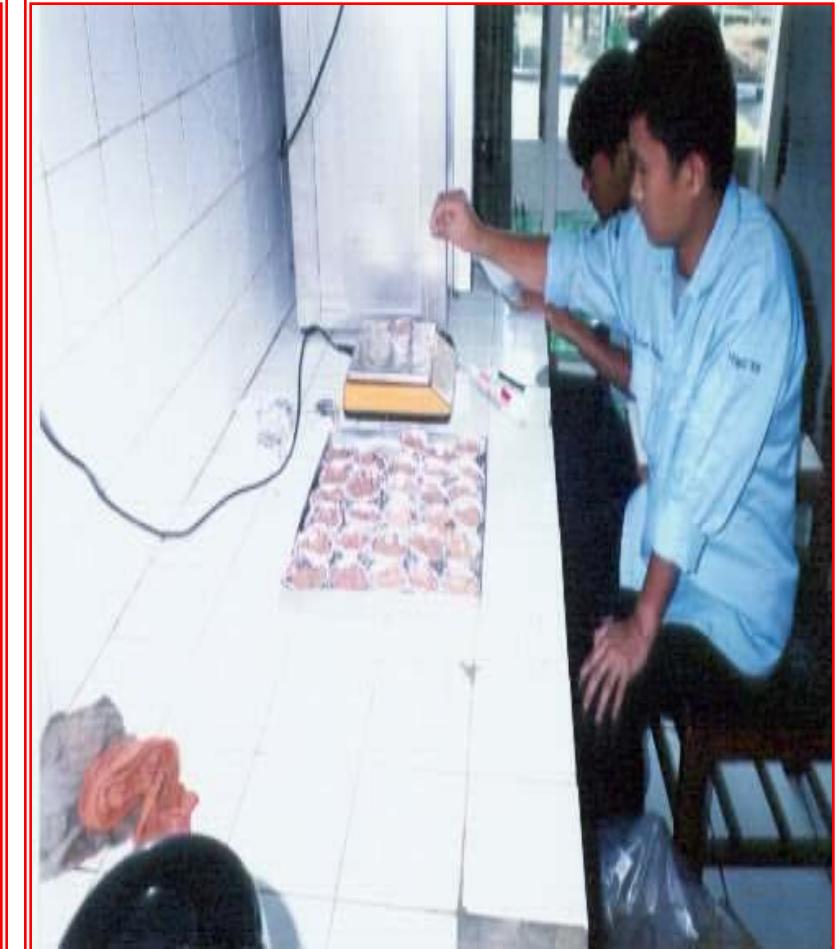
Pengambilan Sampel Tanah Terganggu dan Deskripsi Sifat Tanah di Lapangan



Sampel Tanah untuk Uji Sifat Fisik dan Pengangkutan Sampel ke Lab. Puslittanak Bogor



Analisis Kelembaban Tanah di Lab. BLKKP Lembang



Tanah yang Siap Oven dan Dua Alat Oven yang Digunakan



Hubungan $F(t) = f(R(t))$ Empirik

- Formula Umum :

$$F(t)_{Cr} = 10^{K_{Cr}} - 1$$

$F(t)_{Cr}$ = infiltrasi kumulatif empirik (mm)

K_{Cr} = pola hubungan $F(t)$ dengan $R(t)$

Cr = palawija (W); agroforestri (A); tidak digarap (N); kayu campuran (N); dan pemukiman (P)

- K_{Cr} adalah :

$$K_W = -0,0005 \cdot (R(t)^2 + 0,045 \cdot R(t) + 0,37)$$

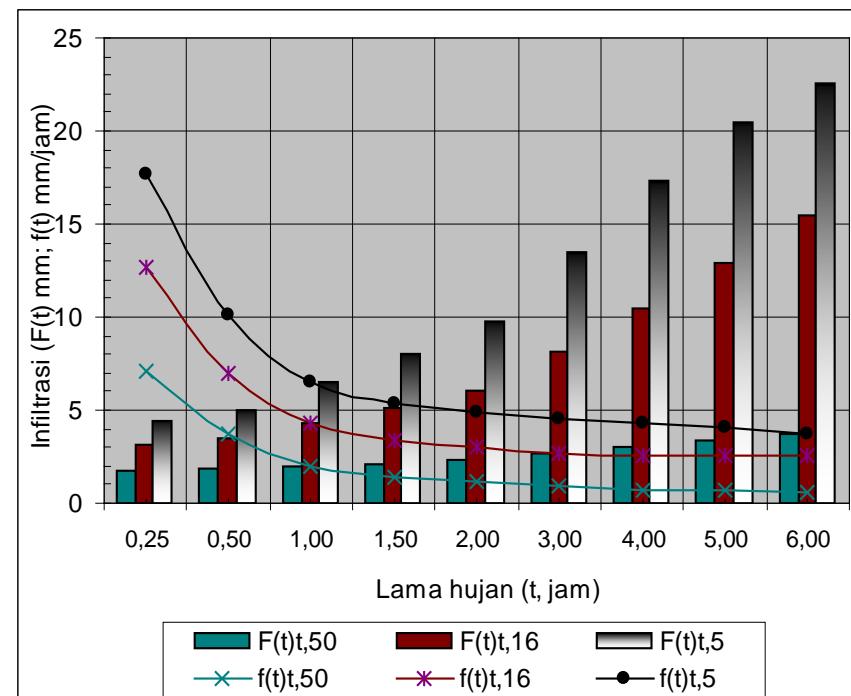
$$K_A = -0,0005 \cdot (R(t))^2 + 0,045 \cdot R(t) + 0,40$$

$$K_N = -0,0004 \cdot (R(t))^2 + 0,039 \cdot R(t) + 0,38$$

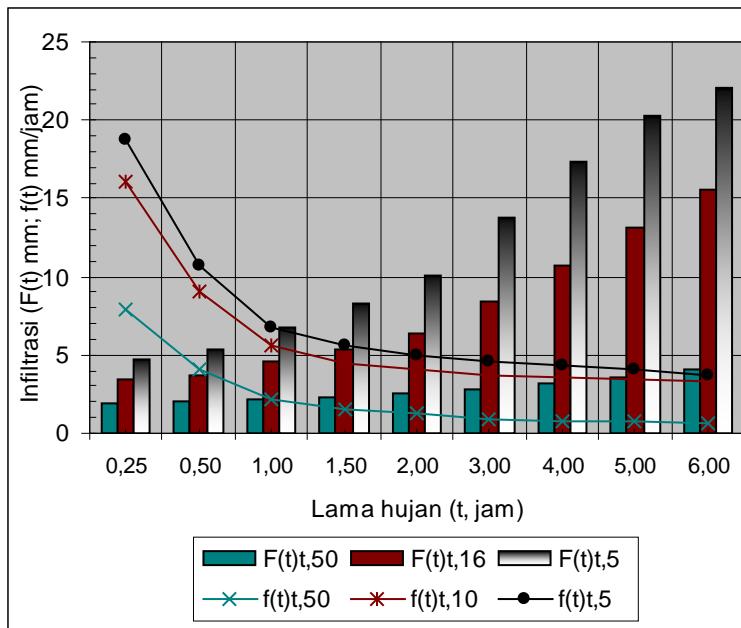
$$K_H = -0,0006 \cdot (R(t))^2 + 0,050 \cdot R(t) + 0,33$$

$$K_P = -0,0004 \cdot (R(t))^2 + 0,040 \cdot R(t) + 0,42$$

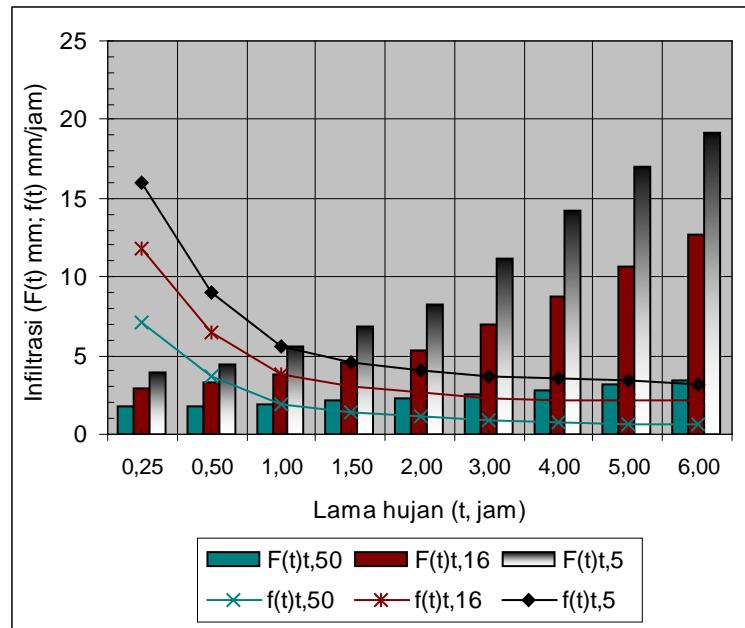
Lahan Palawija



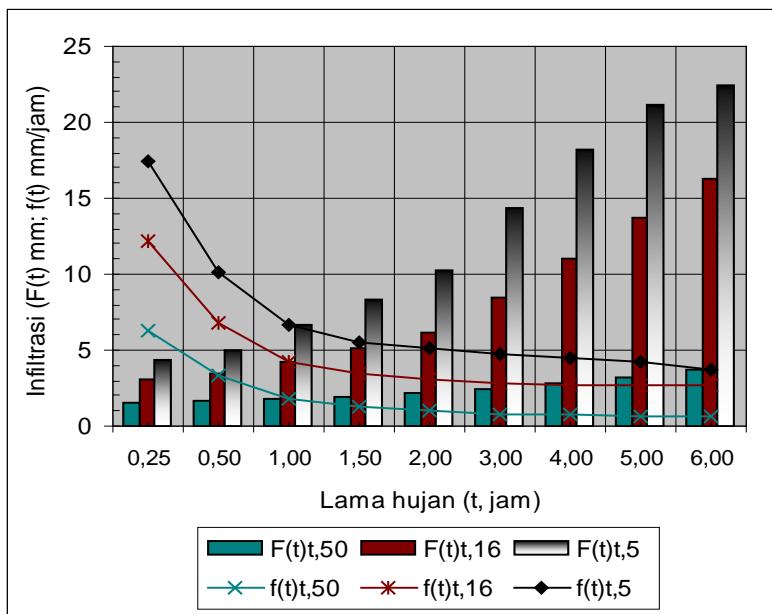
Lahan Agroforestri



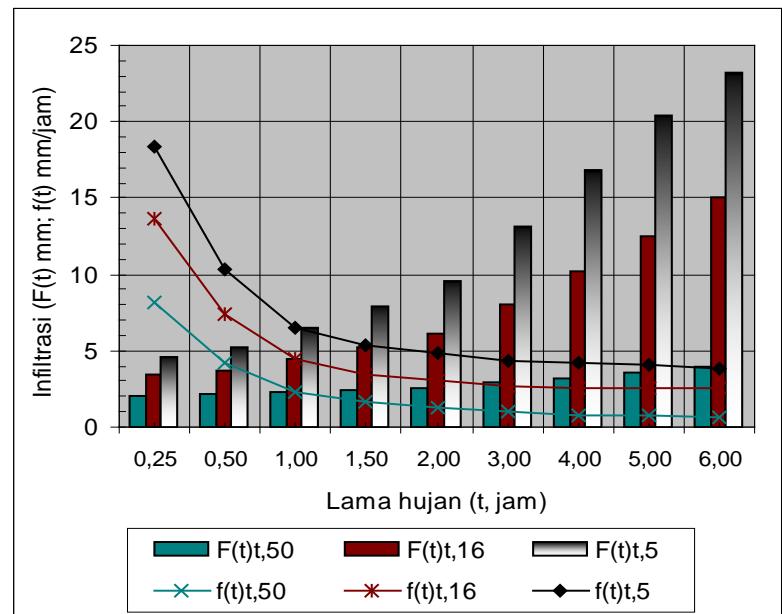
Lahan Tidak Digarap



Lahan Kayu Campuran



Lahan Permukiman



[Data angka...](#)

[Lingkup...](#)

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